



Article

Microbial Communities in Underground Gas Reservoirs Offer Promising Biotechnological Potential

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Abstract: Securing new sources of renewable energy and achieving national self-sufficiency in natural gas have become increasingly important in recent times. The study described in this paper focuses on three geologically diverse underground gas reservoirs (UGS) that are the natural habitat of methane-producing archaea, as well as other microorganisms with which methanogens have various ecological relationships. The objective of this research was to describe the microbial metabolism of methane in these specific anoxic environments during the year. DNA sequencing analyses revealed the presence of different methanogenic communities and their metabolic potential in all sites studied. Hydrogenotrophic *Methanobacterium* sp. prevailed in Lobodice UGS, members of the hydrogenotrophic order *Methanomicrobiales* predominated in Dolní Dunajovice UGS and thermophilic hydrogenotrophic members of the *Methanothermobacter* sp. were prevalent in Tvrdonice UGS. Gas composition and isotope analyses were performed simultaneously. The results suggest that the biotechnological potential of UGS for biomethane production cannot be neglected.

Keywords: methanogenic archaea; methanogenesis; underground gas storage; Power to Methane; green energy; CO₂ utilization

1. Introduction

Microbiological studies of underground gas storage (UGS) have demonstrated the presence of viable microorganisms [1,2] and their significant influence on biological and geochemical processes in these environments [3]. It is likely that Šmigáň and co-workers in 1990 were the first to observe the changes in the composition of the town gas in UGS caused by microorganisms. The recorded decrease in the amount of H₂ and CO₂ combined with the increase in the amount of CH₄ in the Lobodice UGS, Czechia, indicated the

possibility that microbial communities inhabit the UGS. The presence of methanogenic archaea, which appeared to be responsible for the changes in gas quality [3,4], was also confirmed. Thirty years later, our stable isotope measurements confirmed biological methane production at the same site and at other sites. Power-to-gas technology with biological methane production in underground reservoirs seems to be one of the most promising options for carbon-neutral fuel production and storage [5]. The UGS can serve as a reactor where the biological conversion of CO₂ and H₂ to biomethane takes place [6]. Sources of CO₂ can be the thermal sector, various industries, or direct capture of CO₂ from the air [7]. Hydrogen can be produced by electrolysis of water or using surplus renewable energy [8], while hydrogen from waste treatment could also be used in the future [9]. It is also clear that methane will be treated as an energy carrier until safe hydrogen storage and distribution technologies are available [10]. The objective of this study was to determine whether a UGS natural sedimentary rock environment could be a potential target for biomethane production through power-to-methane technology. Although many scientific teams have mapped the microbiome of the subsurface environment (Table 1), the idea of using UGS as a natural bioreactor is new and topical. The enormous capacity of UG enables the production and storage of biological methane in quantities of millions of cubic meters. Three geologically different sites in the Czech Republic—Lobodice, Dolní Dunajovice, and Tvrdonice—serving as UGS were studied in terms of microbial communities using sequencing technology and quantitative PCR to assess their methanogenesis processes.

Table 1. Selected investigations performed in similar rock environment.

Author, Reference	Year	Country	Focus of Study
Šmigán et al. [4]	1989	Czech Republic	methanogenic archaea
Buzek et al. [4]	1994	Czech Republic	microbial methane production
Pedersen et al. [11]	1996	Sweden	microbial diversity
Kotelnikova et al. [12]	1997	Sweden	methanogenic archaea, homoacetogenic bacteria
Fry et al. [13]	1997	USA	microbial diversity
Shimizu et al. [14]	2006	Japan	microbial diversity
Ivanova et al. [2]	2007	Russia	microbiological
Basso et al. [1]	2009	France	microbial diversity
Kimura et al. [15]	2010	Japan	microbial methane study
Flynn et al. [16]	2013	USA	functional microbial diversity
Wu et al. [17]	2016	Sweden	microbial diversity, metabolism
Frank et al. [18]	2016	Russia	variability in microbial composition
Kadnikov et al. [19]	2017	Russia	microbial diversity
Vigneron et al. [20]	2017	USA	microbial methane study

1.1. Methanogenesis

Methanogenesis, the final step in the decomposition of organic matter, is carried out by methanogenic archaea, which play an important role in the global carbon cycle and are responsible for more than half of all methane produced on Earth per year [21]. In the last four decades, methanogenesis has also been described as the dominant metabolic pathway in very deep aquifers [22]. Methanogenic archaea are strictly anaerobic microorganisms and require an environment with low redox potentials of about −300 mV for their growth. Under these conditions, where all other favorable electron acceptors, such as oxygen, nitrate, sulphate, and iron compounds, are depleted or absent, methanogenesis can occur [23].

Methanogenesis is not a uniform process. Methane can be formed via three major pathways: hydrogenotrophic, methylotrophic, or acetoclastic. However, one major characteristic enzyme is present in all types of methanogenesis: methyl-coenzyme M

reductase (MCR), which catalyzes the final step of methyl group reduction to methane [24].

The most widespread and probably the oldest form of methane production is hydrogenotrophic methanogenesis. This pathway is characterized by the conversion of CO₂ and H₂ to methane; molecular hydrogen serves as the electron donor and CO₂ as the electron acceptor. The reduction of formate also occurs via the hydrogenotrophic pathway. The only difference is in the first step, where formate is oxidized to CO₂ and then the CO₂ continues the pathway. The hydrogenotrophic pathway is widely used; for example, by the microorganisms of the orders *Methanobacteriales*, *Methanomicrobiales*, and *Methanococcales*.

The second known pathway is the methylotrophic pathway, in which methylated C₁ compounds (methanol, methylamines, or methyl sulfides) are first activated by specific methyltransferases [25]. Usually, four C₁ compounds are involved in this reaction. One of the methyl groups is oxidized to CO₂ and the remaining three methyl groups are reduced to methane. Methylotrophic methanogens include members of the orders *Methanobacteriales* and *Methanosarcinales* as well as the recently proposed *Bathyarchaeota*, *Verstraetearchaeota*, and *Methanomassiliicoccales*. In addition, some methanogens (members of *Verstraetearchaeota*) were found to be exclusively methylotrophic. Based on this, it was hypothesized that methylotrophic methanogenesis evolved as an independent ancient pathway [26].

Furthermore, two processes have been described for methanogenesis from acetate. The first is acetoclastic methanogenesis, where acetate is cleaved into carboxyl and methyl groups. The substrate for methanogenesis is acetate, which undergoes a disproportionate reaction. The carbon in the methyl group is reduced to methane, while the carbon in the carboxyl group is oxidized to CO₂. Two main genera of methanogens (*Methanosarcina*, *Methanotherix*) are able to use this pathway [27]. The second process is based on syntrophic mutualistic reactions. Less is known about this syntrophic metabolism, which catalyzes the oxidation of acetate to hydrogen and carbon dioxide (SAO) by SAO bacteria. It is possible that syntrophic acetate-oxidizing (SAO) bacteria facilitate acetate consumption and could be coupled with hydrogenotrophic methanogenesis (SAO-HM) [28].

2. Materials and Methods

2.1. Locality and Geological Preconditions

The Lobodice UGS is an aquifer type of UGS located in the central part of the Carpathian Foredeep and formed by an anticlinal structure at an average depth of 450 m. The basement is formed by Proterozoic and Paleozoic crystalline rocks, such as amphibolitic shales. The overlying reservoir sediments consist of lower Badenian clastic sediments, mainly conglomerates, and the caprock sealing the entire structure consists of lower Badenian clays.

The Dolní Dunajovice UGS is a depleted gas reservoir located in the southern part of the Carpathian Foredeep filled with Miocene sediments. Its structure consists of Carpathian sediments, mudstones, sandstones, and siltstones overlying the Eggenburgian sediments. The Eggenburgian mudstones form the caprock of the UGS and the Eggenburgian basal clastics glauconitic sandstones serve as the UGS horizon at an average depth of 1100 m. Jurassic carbonates (Kurdejov limestones) form the UGS basement.

The Tvrdonice UGS is a depleted set of gas and oil reservoirs located in the Czech part of Vienna basin. The sedimentary complex overlies the crystalline rocks of Brunovistulikum, and the thickness of the sediments reaches over 5 km in some parts of the Vienna basin. The hydrocarbon reservoirs, later converted to UGS, are located at a depth of 900 to 1600 m from the Sarmatian–Badenian boundary interval.

Gas and oil fields are generally structural traps tied to a fault system and subdivided into specific smaller segments and horizons. The reservoir horizons are formed by sandy strata interbedded with mudstones.

2.2. Sampling

Two types of water samples were collected from 16 wells over a two-year period. Water samples from overflow and groundwater were collected using a special rope technique and a sterile subsurface sampler—a stainless steel capsule with a volume of 0.75 L (Leutert GmbH, Adendorf, Germany).

Samples were collected from the overflow by passing water from the well through a 200 L barrel into which a sterile 1 L glass bottle was submerged. This sample was used for molecular biological analyses. For sample cultivation, a 5 L canister with a drain valve was filled in a similar manner. The water sample (30 mL) from this canister was then transferred to a sterile 150 mL culture vial fitted with a rubber stopper. A sterile tube was connected to the drain valve of the canister, at the end of which a sterile needle was inserted into a Luer-Lock thread to pierce the rubber stopper of the vial. The sterile culture vials were flushed with nitrogen several times before use. The nitrogen was then aspirated, and the vial gas phase was under vacuum so that the sample could be aspirated from the canister. For the subsurface sampler, we connected a sterile hose to the sampler valve after releasing the pressure in the sampler, which led to a 0.5 L sampling tube. As the water flowed through the tube, we punctured the wall of the tube with a sterile syringe and needle and collected water for culture. This water (30 mL) was transferred with a syringe to a sterile culture vessel that was under negative pressure. The water that flowed through the tube into the sterile 0.5 L sample tube was then used for molecular biology analyses. In both cases, culture was performed according to the procedures described in [29,30]. Water samples were transported to the laboratory under anaerobic conditions in a special cooling box. A total of 20 samples were collected, comprising 11 overflow water samples and 9 groundwater samples (Table S1).

2.3. Physical-Chemical Parameters and Groundwater Chemical Composition

Measurements of pH, redox potential (ORP), electrical conductivity, and temperature were conducted on site using the WTW Multi 350i (accuracy ± 0.01 for pH, ± 0.2 for ORP, $\pm 0.5\%$ for electrical conductivity and ± 0.1 °C for temperature). The SenTix 41 electrode was used to measure pH and temperature. The SenTix ORP electrode was used to measure redox potential. The ORP value measured against Ag/AgCl electrode was recalculated against standard hydrogen electrode (SHE) according to the operating manual. The WTW TetraCon 325 electrode was used to measure electrical conductivity. Prior to analysis, all groundwater samples were filtered through a 0.45 μm membrane filter (Millipore, HAWG047S6). Chemical analyses of the main groundwater components were performed in the chemical laboratory of the Department of Geological Sciences, Faculty of Science, Masaryk University, according to standard laboratory procedures. Basic data processing was performed in Microsoft® Excel spreadsheet, advanced data processing, geochemical calculations and geochemical modelling were performed in Geochemist's Workbench®, release 12.0.4. (Aqueous Solutions LLC, Champaign, USA).

2.4. Degas Analysis

2.4.1. Isotopic Determination

Isotopic determination of $\delta^{13}\text{C}$ in CH_4 and δD in CH_4 of the gas samples dissolved in water was performed using a Picarro Cavity Ringdown Spectrometer (CRDS), a G2201-*i* Analyzer for Isotopic CO_2/CH_4 , and a G2182-*i* Analyzer for δD & $\delta^{13}\text{C}$ in CH_4 . Isotopic determination of $\delta^{18}\text{O}$, δD , and $\delta^{17}\text{O}$ in water was performed using the CRDS L2140-*i* Analyzer for Isotopic H_2O (all by Picarro, Inc., Santa Clara, USA). Instrument setup and sample preparation in terms of appropriate concentration range, standards, etc. are described in the manufacturer's guidelines and in the standard operating procedures of Testlab Geo-Services (RWE Gas Storage CZ, Ltd., Brno, Czech Republic).

2.4.2. Gas Chromatography

Analysis of the gas samples was performed using the Agilent 7890B Gas Chromatograph (Agilent Technologies, Inc., Santa Clara, USA), a three-channel system using TCD-TCD-FID detectors. The gas chromatograph was equipped with two precolumns HiSep Q, and three separation columns (HP-Plot Q, Molsieve 5a, and HP Molsieve), all of which were 0.53 mm in diameter. The mobile phase was Ar 5.0 and He 5.0 (SIAD Czech, Ltd., Rajhradice, Czech Republic). The thermal program and other parameters were set according to the standard operating procedure of Testlab Geoservices (RWE Gas Storage CZ, Ltd., Brno, Czech Republic). (Results shown in Table S2).

2.5. Microscopy

Samples selected for analysis by Scanning Electron Microscopy (SEM) were prepared to withstand the high vacuum conditions in the SEM chamber. An appropriate volume (20 mL) of the water sample was filtered onto a polycarbonate membrane filter (0.2 µm, Merck Millipore, Guyancourt, France) using a vacuum filtration device (Merck Millipore, Guyancourt, France). The cells on filters were fixed with 2% glutaraldehyde in 0.1 M sodium cacodylate buffer (pH 7.4). After one hour at room temperature, followed by overnight fixation and drain off fixative, the samples were immediately transferred to 50% ethanol solution (EtOH). After fixation, small pieces of the filters were dehydrated by a graded EtOH series (70%, 85%, 95% and twice 100% EtOH), each step taking approximately 20 min at room temperature. The filter pieces saturated with 100% EtOH were dried with CO₂ (K850 Critical Point Dryer, Quorum Technologies, Lewes, UK) at the critical point and placed on a stub with conductive carbon tape. To increase the conductivity of the samples, the filters were sputter-coated with 5 nm of iridium (Q150T ES, Quorum Technologies, Lewes, UK). Samples were analyzed using a high-resolution field emission scanning electron microscope TESCAN CLARA (TESCAN ORSAY HOLDING, Brno, Czech Republic). All images were acquired at a low accelerating voltage of 1 keV using the in-column Axial detector.

2.6. Molecular Biological Methods

2.6.1. DNA Isolation

Well water samples (0.5 L from the subsampler, 1.5–2 L from the overflow) were filtered on 0.22 µm membrane filters (GTTP, Millipore, France) and subjected to DNA isolation using a kit according to the manufacturer's instructions (DNeasy Power Water Kit, Quiagen, Hilden, Germany). DNA quality was measured using NanoDrop 2000 UV-Vis spectrophotometer and Qubit TM fluorometer (Thermo Fisher Scientific, Waltham, USA).

2.6.2. Quantitative PCR

The qPCR (quantitative polymerase chain reaction) method was used for absolute quantification of methanogens. A pure culture of *Methanobrevibacter smithii* was used as a template for the standard. DNA concentration was measured using a fluorometer Qubit (Thermo Fisher Scientific, Waltham, USA) and the sample was diluted to the required concentrations (10⁴–10⁸ copies per µL). The reaction was performed on Light Cycler 480 (Roche, Switzerland) in triplicate for each sample. The reaction volume was 14 µL, including 4 µL of template DNA and 9 µL of Luna Master Mix (BioLabs, Ipswich, USA) with two forward (0.25 µL per one) and one reverse primer (0.5 µL), each at a final concentration of 250 nM. The primer was targeted to the *mcrA* gene, which is supposed to be a single copy gene (one gene per methanogen cell). A combination of three primers was designed for this study. The reverse *mcrA* primer 5'-CGTTCATBGCAGTAGTTVGGRTAGT-3' was used in combination with equal volumes of the two forward primers *mcrAF1* 5'-ACTTCGGTGGATCDCARAGRGC-3' and *mcrAF2* 5'-ACTTCGGCGGTTCDCARAGRGC-3' [31,32]. The reaction conditions included an initial denaturation step at 95 °C for 3 min, followed by 35 cycles of denaturation at 95 °C

for 30 s, annealing at 55 °C for 45 s, and extension at 72 °C for 30 s, with a ramp rate of 0.1 °C/s from the annealing to the extension temperature, followed by a final extension step at 72 °C for 10 min. Due to fluctuations in the signal during the first five cycles, fluorescence was read from the sixth cycle onward. The expected length of the amplicons was approximately 300 bp.

2.6.3. Illumina—Next Generation Sequencing Method

DNA extracted from groundwater was used as the template for PCRs with specific primers flanking the V4 region of the 16S rRNA gene sequence [33]. Amplification was performed using Platinum™ II Taq Hot-Start DNA Polymerase (ThermoFisher) at 0.8× according to the Earth Microbiome protocol [34]. After PCR, the amplification products were purified using Agencourt® AMPure XP Beads (Beckman Coulter, Brea, USA). Subsequently, the purified PCR samples were quantified and normalized using the Qubit fluorometer. The normalized PCR products were pooled, and their length and quality were checked using the DNF-474 HS NGS Fragment Kit for Fragment Analyzer (Agilent, Santa Clara, USA). The final library was sequenced using an Illumina MiniSeq sequencer together with the Mid Output Kit (2 × 150 paired end sequencing) according to the manufacturer's instructions (Table S3).

Raw fastq reads were processed using the DADA2 package (version 1.16.0), [35] in R (version 4.0.0). Analysis was performed according to the standard operating procedure. Reads were first filtered, then trimmed, de-replicated, and de-noised. Then, forward and reverse reads were merged, chimeras were removed, and the taxonomy was assigned by the RDP naive Bayesian classification [36] against the Silva database [37]. Multiple alignments were performed using the DECIPHER package and a phylogenetic tree was constructed using the phangorn package [38]. Phylogenetic and statistical analyses were then performed in R using the phyloseq package [39]. The datasets generated and analyzed during the current study are available in SRA under the project number BioProject ID: PRJNA759841.

To assess the metabolic potential of the communities based on the 16S rRNA gene amplicon data alone, functional annotation was performed using the FAPROTAX database [40]. The normalized and curated OTU abundances were assigned to a phylogenetically conserved functional group from information based on functional annotations of cultivated representatives. The program assumes that all cultivated and non-cultivated members can perform the functions verified in the database. The total DNA extracted from UGS water samples does not exclusively represent the metabolically active part of the community, as DNA from dormant and dead organisms is also extracted simultaneously [41]. For a deeper understanding of the metabolism of the microbial community metabolism, transcriptomic or proteomic data would be required and should be considered.

3. Results

3.1. Physicochemical Parameters and Water Chemistry

Characteristics of Groundwater

The chemical composition and physicochemical parameters of the studied UGS groundwaters are shown in the diagrams of Piper and Durov in Figure S1. Lobodice UGS groundwaters are slightly alkaline, pH is in the range of 7.36–8.70, and redox potential varies in the range of –238 to –399 mV, which means strongly reducing conditions. The groundwater has a relatively low mineralization of 2.4–3.1 g/L. Among the cations, sodium ions are predominant; among the anions, chlorides and bicarbonate are comparable, with bicarbonate slightly predominant. Water from caprock shows about twice as much mineralization. The ratio of major anion and cation concentrations is almost the same as for typical groundwaters from the Lobodice UGS. The groundwaters from the Tvrdonice UGS are also slightly alkaline (pH 8.10 and 8.45) with reducing conditions (–130

and -115 mV). Compared to the Lobodice UGS, they have about three times the mineral content (10.21 and 10.61 g/L). Their chemical composition is quite simple: of the main constituents, only sodium and chloride ions are present, while the concentrations of the other constituents are only in the two-digit milligram range. The groundwater of the Dunajovice UGS is almost identical in chemical composition to the groundwater of the Tvrdonice UGS, except that it has a three-times higher mineral content of about 30 g/L. The pH values are also in the characteristic range (7.65 and 8.07), and the redox potential indicates a typical anoxic environment (-138 and -142 mV).

All three types of UGS groundwaters are typical synsedimentary waters, and their properties reflect the sedimentary conditions and characteristics of the rock environment in which they are contained. The carbon, nitrogen, and sulfur components are major constituents that determine the oxidation–reduction processes. For the carbon component, the pH values are just around the groundwater saturation limit for calcite (Figure S2a), indicating that the groundwater is saturated relative to calcite. The solubility product of calcite determines the concentrations of calcium and bicarbonate ions, and conversely, pH values are buffered by the equilibrium between dissolved carbonate species and calcite.

The redox potential of the samples where groundwater was discharged from the wells for several hours is, with two exceptions, in the range where nitrogen species are stable as NH_4^+ (Figure S2b), and approach the limit of redox transformation of $\text{SO}_4^{2-}/\text{HS}^-$ —manifested by pyrite precipitation under the given conditions (Figure S2c). Pyrite is present in the UGS sedimentary rocks and provides reducing conditions. As for carbon species, the pH and Eh conditions are close to the transition boundary between dissolved carbonate species (CO_2 and HCO_3^-) and methane (CH_4). The speciation diagram for carbon species (Figure S2d) shows that the conversion of carbon dioxide to methane begins at a redox potential of -225 mV. The physicochemical conditions and chemical composition of the studied UGS groundwaters are suitable for biologically assisted methanation, i.e., the conversion of carbon dioxide to methane by the action of molecular hydrogen. These conditions are supported by other redox-active components (NH_4^+ and HS^-).

3.2. Isotopic Determination

The results of the isotopic analysis of the gas and water samples indicate the origin of these samples. These results and an example of the typical isotopic composition of the injected gas are given in Table S4. The raw δD values measured in CH_4 must be corrected using the δD values measured in water, because there is an equilibrium between these values, and the values in methane are affected by the values in water. The data listed in Table S4 are also shown in Figure 1. The structure in the background [42] helps to immediately assign the samples to a specific type of origin. As can be seen in Figure 1, the samples from Lobodice and Dolní Dunajovice are clearly of microbial origin. The results for the samples from Tvrdonice indicate that these degassing samples are a mixture of microbial and thermogenic gas. For comparison, the values for injected gas have also been added to this figure to illustrate the difference between the thermogenic gas typically stored in the UGS and the gases that were influenced (at least to some degree) by microbial processes.

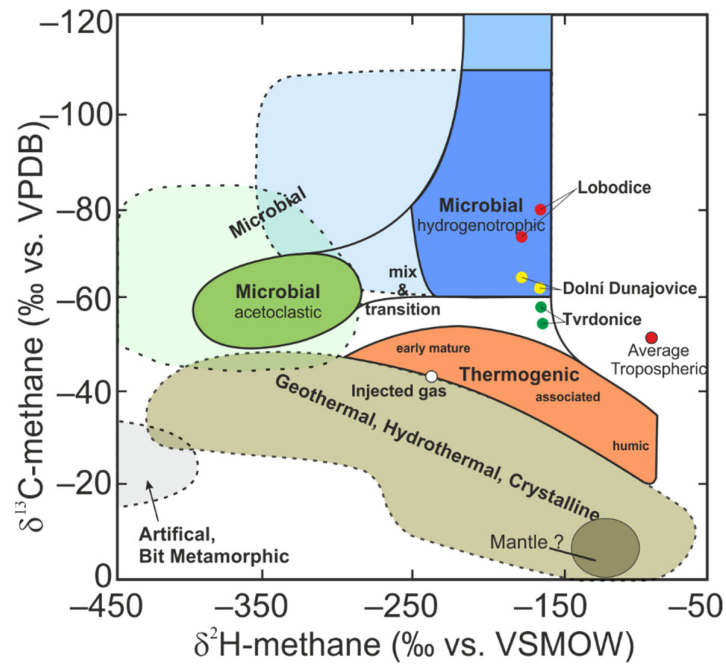


Figure 1. C–D plot [42] with measured data.

3.3. Microscopy

The SEM-micrograph shows a wide variety of morphological types of microorganisms, predominantly the accumulation of rod-shaped cells with a length of 1–4 μm and cocci with a diameter of about 0.3 μm (Figure 2).

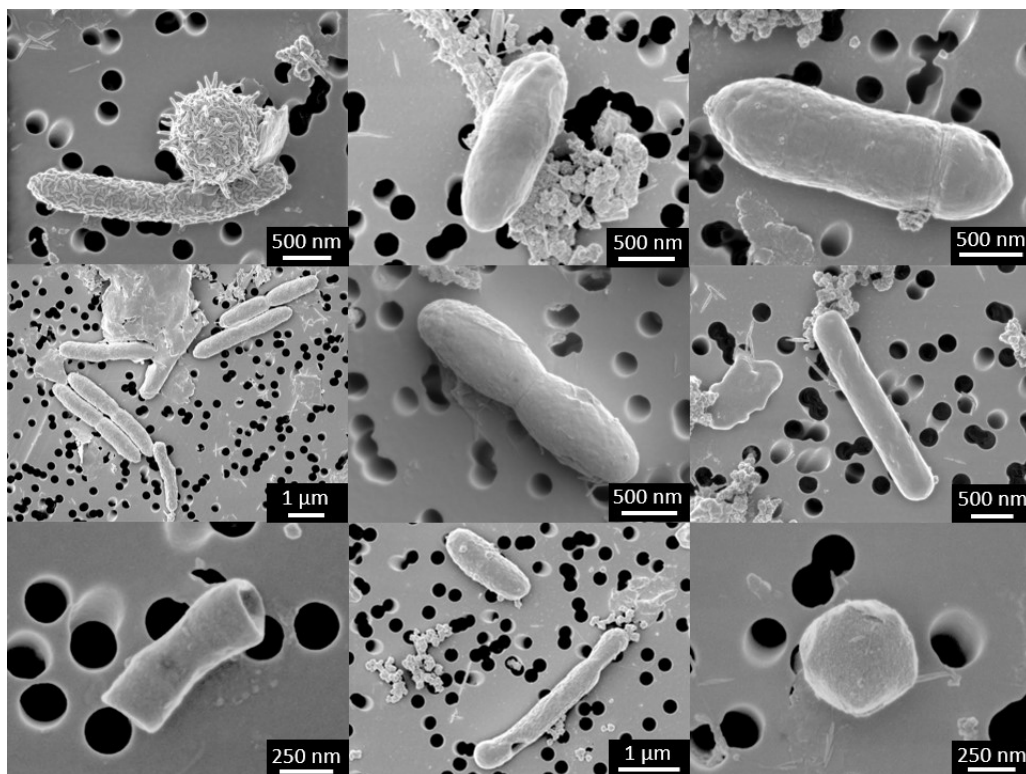


Figure 2. Examples of microbial morphotypes: rods and cocci found in well TVR-B captured by SEM. Inorganic nanoparticles which form conglomerates can be seen.

3.4. Quantitative Analyses of Methanogens (qPCR)

Of all sites, 14 wells were sampled, ten of which were sampled once and four of which were sampled repeatedly in different seasons, focusing on specific UGS regimes (injection–extraction gas periods) (Table S2). Of the total twenty samples from the dataset, DUN-B (December 2018), LO-H (November 2019), LO-H (October 2018), and LO-I (March 2019) were discarded due to low DNA concentration. The number of *mcrA* gene copies (Table S5) found in the Lobodice samples ranged from $7.20 \cdot 10^1$ to $2.40 \cdot 10^7$ in 1 μL of DNA isolated from 1 L of well water. The highest number of *mcrA* gene copies ($2.40 \cdot 10^7$) was detected in the well LO-C (October 2018), while the lowest number ($2.00 \cdot 10^2$) was found in LO-F (May 2018). In Dolní Dunajovice, the *mcrA* gene copies ranged from $1.7 \cdot 10^5$ to $1.03 \cdot 10^6$, and in Tvrdonice from $1.86 \cdot 10^4$ to $1.42 \cdot 10^6$. Certain wells in Lobodice were repeatedly sampled at different times of the year over a two-year period. When comparing the number of *mcrA* gene copies, the hypothesis about the influence of the UGS regime (injection/withdrawal periods for gas) on methanogen abundance cannot be confirmed.

3.5. Metagenomic Analyses of UGS Archaeal Community

3.5.1. Next-Generation Sequencing (NGS) Analysis

Next-generation sequencing (NGS) analysis targeting archaeal and bacterial 16S rRNA genes was performed to elucidate microbial community structures in anoxic groundwater from deep aquifers. The 24 samples were sequenced along with 16 independent samples on Illumina MiniSeq using the MidOutput Kit (2×150 bp). The total run yield was 2.73 Gbp, with 88.39% of the reads passing the quality filter ($>Q30$), resulting in 8,689,550 reads passing the filter. The average error rate of the sequencing run was 0.98%. The DADA2 algorithm extracted 2864 unique ASVs from the 24 samples (Table S6). The sequencing depth obtained far exceeded the requirements, as shown by the rarefaction curve (Figure S3) with a minimum of 94,616 reads and a mean of 150,555 reads per sample.

For Illumina MiniSeq sequencing, the V4 region of the 16S rRNA gene was amplified [34,43]. The 16S rRNA gene is ubiquitous, occurs in several bacterial and archaeal species, and is highly conserved. Members of both domains were detected in varying unexpected proportions in all samples (archaea comprised 0.2%–75.7% of the microbial community in all sampled wells). Archaeal community composition differed slightly among the three sites, as did environmental and physicochemical conditions (Figure S4).

All samples collected were positive for the presence of methanogenic archaea. The 16S rRNA marker was used for sequence analyses and indicated the presence of an archaeal community in each well (Figure 3). The archaea were represented by five identified phyla, *Euryarchaeota* and *Crenarchaeota*, *Hydrothermarchaeota*, *Nanoarchaeaeota* and *Hadesarchaeaeota*, the last being a recently proposed phylum of thermophilic microorganisms found in deep mines, hot springs, marine sediments, and other subsurface environments [44]. This includes eight discovered classes, seven orders, thirteen families and fifteen genera. Sequencing of the 16S rRNA gene shows that the majority of the 140 archaeal OTUs belong to methanogens and consist of 10 genera and two recently discovered *Candidatus* species, whose abundance accounted for more than 5% of the total community.

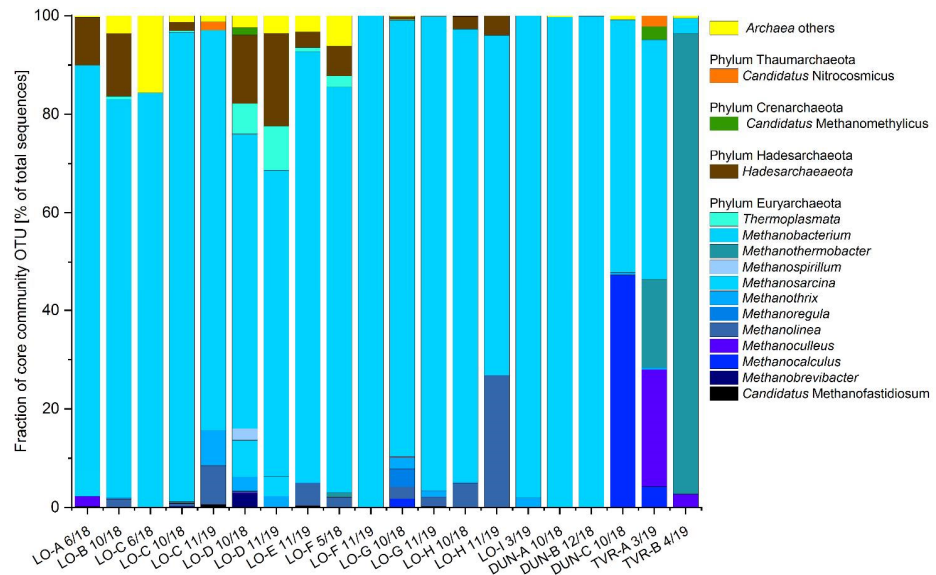


Figure 3. Archaeal community—16S rRNA gene sequence focused on methanogens detected in water samples from UGS’s (5% cut).

The composition of the methanogenic archaea community detected in the UGS water samples shows that the major metabolic pathway in the UGS environment is hydrogenotrophic methanogenesis, with hydrogen and carbon dioxide as carbon and energy sources (Figure 4). In general, the genus *Methanobacterium* predominated in almost all samples, followed by the genus *Methanothermobacter* (present exclusively in Tvrdonice) and *Methanolinea*. The metabolism of these genera is exclusively hydrogenotrophic methanogenesis. The acetoclastic members of the genera *Methanotrinx* and *Methanosarcina* were represented by large numbers of cells in the LO-A and LO-D wells.

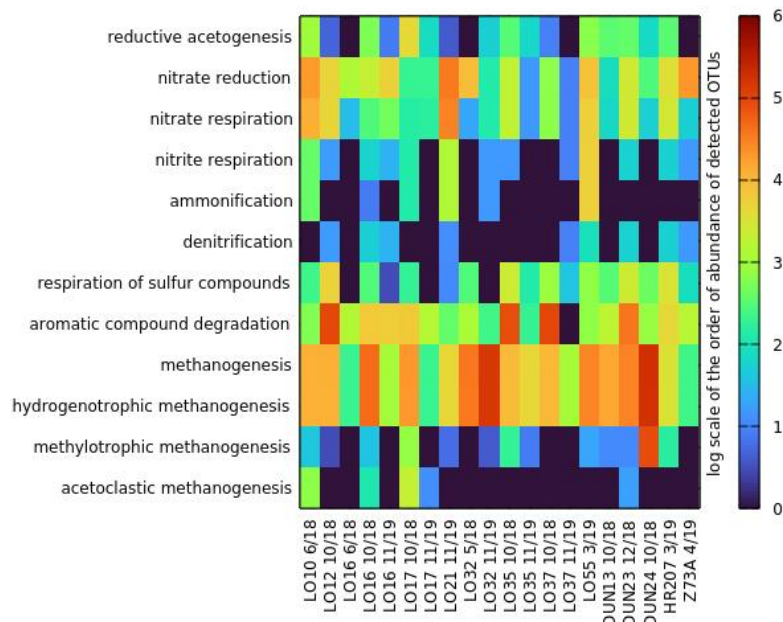


Figure 4. Selected metabolic groups predicted using the database of functional annotations of prokaryotic organisms, FAPROTAX.

Exceptionally high numbers of members of the genus *Methanocalculus* were found in the DUN-C well, where they make up half of the archaeal community. *Methanothermobacter* genera absolutely dominate in TVR-B and contribute to the archaeome in TVR-A, as do *Methanoculleus* members. In addition, new methanogens of interest were detected in interesting numbers in this well, namely the methylotrophs *Candidatus Methanomethylicus* and *Candidatus Methanofastidiosum*.

3.5.2. Biodiversity of Microbial Communities

Shannon and Unifrac indexes were chosen to describe the biodiversity of microbial communities. Alpha diversity refers to the diversity within a given ecosystem and is usually expressed by the number of species (or species richness). The diversity and richness of microbial communities inhabiting the different wells were determined using 16S rRNA gene analyses. Comparing the microbial diversity of each sample, Lobodice recorded the highest values with respect to all three sampling sites. Within all localities, the highest alpha diversity was found in the sample LO-D 11/19, while the opposite was found in LO-C 10/18. Relatively high values for the Shannon index were found in LO-C 11/19, LO-F 5/18, LO-F 11/19, and DUN-A 10/18 (Figure S5, Alpha biodiversity measure).

Beta diversity describes the structural complexity of the environment. It is a measure of the difference (or conversely, similarity) in species composition between communities along a given gradient of the environment, or between the community and its environment. Beta diversity is higher when a community contains uncommon species. Beta diversity shows the difference between microbial communities from different sampling sites. The obtained indices show no significant clustering among all samples. Only one distinct group formed from samples from Dolní Dunajovice and Tvrdonice, which originated from similar environments. In addition, some samples from Lobodice (LO-D 10/18, LO-E 11/19, LO-A 6/18) were clustered together with them. The most divergent samples, considering all sites, were LO-C 10/18 and LO-G 10/18 (Figures S5 and S6).

3.5.3. Metabolism Prediction—FAPROTAX

The potential metabolic functions of microorganisms were estimated using the database of functional annotations of prokaryotic organisms, FAPROTAX, which showed that the most frequent categories were related to carbon cycling under anaerobic conditions. Furthermore, metabolism involved in nitrate respiration and respiration of sulfur compounds was found to be marginal (Figure 4). During anaerobic decomposition of organic material, the hydrolysis products are decomposed into simple organic and inorganic substances (acids, alcohols, CO₂, H₂). Fermentation of these substances produces several reduced end products. In the next step—acetogenesis—the syntrophic acetogenic microorganisms produce hydrogen and decompose organic acids, alcohols, and some aromatic compounds. In the dataset, aromatic compound degraders, such as members of the genera *Pelotomaculum* or *Acinetobacter*, were detected in all samples, but significant levels were reached in the wells LO-B, LO-G, LO-H, and DUN-B. The presence of syntrophic acetogenic bacteria (*Sporomusa*, *Anoxyratronum*) was confirmed in the LO-D well. The dataset also reflects the composition of the methanogenic community in the well, showing a high abundance of acetoclastic methanogens. Acetogenic microorganisms are commonly found in syntrophy with methanogenic archaea, which consume the hydrogen they produce. Methanogens utilize substrates, such as carbonaceous substances (methanol, formic acid, methylamines, CO₂, CO) or acetic acid. The end products of their metabolism are methane and carbon dioxide. All types of methanogenic metabolism were recorded in the collected samples, but hydrogenotrophic methanogenesis, represented mainly by the genus *Methanobacterium*, was predominant. A significant potential of methylotrophic methanogenesis was observed only in DUN-C (*Methanocalculus*, *Methanospirillum*). Acetotrophic methanogens, represented exclusively by the genus *Methanosaeta*, were detected only in the LO-D well, corresponding to high acetogenic

activity. The highest methanogenic activity was then found in DUN-C and LO-F. The intensity of predicted methanogenic metabolism varied by site and sampling, showing unexpected changes in metabolic activity. In general, it appears that this very particular type of ecosystem is more dynamic than should be expected.

4. Discussion

Over the course of four years of studying biomethanation in underground water reservoirs, we have built on the results summarized in the review paper [5].

The objective of the study was to determine part of the composition of the UGS microbiome, focusing on methanogenic archaea, biological methane production, and evaluation of the biotechnological potential for biomethane production from underground hydrocarbon reservoirs and aquifers. Although methanogens have been detected in aquifers [4,18,19] or in the sediments of UGS pipelines [45], to the best of our knowledge, this is the first thorough study addressing different types of reservoirs. We hypothesize that seasonal fluctuations of gas in UGS may have an impact on the abundance of methanogens in UGS [2].

To demonstrate the dynamics of their abundance during the year, some of the wells were sampled at different times of the year, focusing on specific UGS regimes (injection/withdrawal gas). The results show that the methanogenic community is dynamic throughout the year, but no significant trend was found with respect to the UGS regime (Figure 3). The results of the sequencing analyses confirmed that each well hosts a specific methanogenic community that accounts for 70%–100% of the total archaeal community. This means that methanogen abundance in UGS is generally greater than 10%.

Hydrogenotrophic *Methanobacterium* sp. prevailed in most samples from the Lobodice locality, as proven by NGS and cultivation. *Methanosarcina* sp. were also abundant, consuming acetate or methylated compounds as a substrate for methanogenesis. The second sampling site, Dolní Dunajovice, had a different reservoir environment and thus a different microbiome composition compared to Lobodice. Members of the hydrogenotrophic order *Methanomicrobiales* were most abundant. The third sampling site, Tvrdonice, was strongly influenced by the higher temperature in the UGS and therefore contained thermophilic members of the *Methanothermobacter* genus.

Methanogenic and fermentative microorganisms are often organized in mutualistic consortia to facilitate rapid electron exchange by diffusion of hydrogen or formate [46]. In addition, electron exchange by direct electron transfer between species has been discovered. *Geobacter metallireducens* transfers electrons directly to *Methanothermobacter* during methanogenic degradation of ethanol, presumably through nanowires [47]. Members of the genera *Geobacter* and *Methanothermobacter* were found in samples from LO-A and DUN-A.

Anaerobic syntrophy is defined as a thermodynamically interdependent extreme lifestyle in which the degradation of an organic compound occurs only when the end products (usually hydrogen, formate, and acetate) are maintained at very low concentrations. Microbial syntrophy between *Bacteria* and methanogenic archaea enhances the methanogenic activity and methane yield. This type of syntrophy is related to the global carbon cycle in anaerobic environments, which is based on a complex community of metabolically coupled microorganisms that are highly adapted to the environment. This was clearly demonstrated in the publication [48].

The dynamics of the conversion of H₂ and CO₂ to methane varied in all samples depending on the UGS environment. The key parameters were temperature and chemical composition of the UGS groundwater. Biomethane production is also influenced by the composition of the microbial community in each UGS well [49]. Our study confirms the presence and natural activity of methanogenic archaea in underground gas storages. The function of underground gas storages as natural bioreactors is confirmed by the result of our field experiment under real reservoir conditions [50]. The strategic importance of this

solution is quite clear. These principles can make an important contribution to reducing the impact of transport and energy on nature and decarbonizing the economy. The use of underground reservoirs for industrial production of biological methane is one of the ways by which the Czech Republic might achieve self-sufficiency, avoiding dependence on natural gas imports.

4.1. Lobodice

The low salinity and higher pH (close to eight) of the environment provide ideal conditions for the growth of members of the genus *Methanobacterium*. The high prevalence of this genus was confirmed in all but one of the sampled wells. This genus is hydrogenotrophic but can also metabolize formate. Several species have already been isolated from aquifers, so these genera seem to be widespread in this environment [3,15,51,52]. Moreover, this genus has even been isolated during previous experiments at the Lobodice UGS [4]. Members of the genus *Methanosarcina* were found in large numbers in LO-A and LO-D. The main substrates for growth are acetate, methanol, trimethylamines, or other methyl-containing compounds. Some of the species can form methane from H₂ and CO₂, or they use H₂ to reduce methanol to methane. Acetotrophic members of the genus *Methanotherix* were detected and were present in samples depending on morphological observation, especially in LO-C, LO-D, and LO-G samples. They are exclusively served by acetate as a substrate, and their affinity for this substrate is much higher than that of the genus *Methanosarcina*. Members of the genus *Methanoculleus* were found exclusively in the well LO-A. This genus usually lives in marine environments and brackish water, but is also widespread in other environments, such as bioreactors, landfills, or wastewater. Unlike other Archaea, *Methanoculleus* can use ethanol and some secondary alcohols as electron donors for final methane production [53]. Conditions at this site appear to support acetogenic bacteria of the genus *Acetobacterium*, which use H₂ and CO₂ to form acetate. The acetate formed does not remain in the liquid for long, being rapidly consumed by a narrow range of bacteria, or serving as a substrate for methanogenesis.

4.2. Dolní Dunajovice

The Dolní Dunajovice site is characterized by higher salinity and temperature compared to Lobodice. Due to these conditions, the composition of methanogens differs significantly from that in Lobodice. In the wells DUN-A and DUN-B, methanogens belonging to the order *Methanobacteriales* absolutely dominate. In the sample from the well DUN-C, two predominant archaeal taxa, *Methanobacterium* sp. and *Methanocalculus* sp., comprised around half of the organisms. Species of the genus *Methanocalculus* are very salt tolerant and can live at sodium chloride concentrations as high as 125 g/L [54]. Moreover, the higher temperature in the Dolní Dunajovice reservoir is ideal for them, as the optimal temperature of the species is 45 °C [55].

4.3. Tvrdonice

Tvrdonice offers unique conditions, with salinity levels somewhere between those of Lobodice and Dolní Dunajovice, at temperatures around 50 °C. The higher temperature fosters microbial communities distinct from those of mesophilic environments. The dominant methanogen, *Methanothermobacter* sp., was found in well TVR-B and was also detected in another deep aquifer [15]. Its predominance is likely to lead to faster reactions in methane formation. High temperatures are a key factor affecting microbial composition. The absolute prevalence of the thermophilic *Methanothermobacter* sp. (94%), which grows best at temperatures between 55 °C and 65 °C, was confirmed by 16S rRNA sequencing. Only two other genera were detected in well TVR-B by 16S rRNA sequencing: *Methanobacterium* (3.2%) and *Methanoculleus* (2.7%). These genera use carbon dioxide and hydrogen as substrates to produce methane for energy production.

Samples from TVR-A well water consisted of the genera *Methanobacterium* (approximately 50%), *Methanococcus* (24%), *Methanothermobacter* (18%), and members of the recently proposed taxa *Candidatus Methanomethylicus* and *Candidatus Methanofastidiosum* [56]. *Candidatus Methanofastidiosum* is a unique methanogen that utilizes methylated thiol reduction and bridged carbon and sulfur cycles and may compete with CO₂-reducing methanogens and even sulfate reducers [56,57].

5. Conclusions

Based on our results, we can conclude that the underground gas storages assessed by us showed a microbiome composition suitable for biological methane production. It can be concluded that if underground gas storages offer suitable habitat, they can be used as fermenters for the biological conversion of CO₂ and H₂ into biomethane anywhere in the world. From a strategic perspective, underground storages can play an important role in the energy mix, as they can be used for long-term energy storage.

Supplementary Materials: The following supporting information can be downloaded at: www.mdpi.com/article/10.3390/fermentation8060251/s1-s6, Table S1: Overview on sampling and on physiochemical parameters measured in sampled wells, Table S2: Table of gas composition measured in sampled wells, Table S3: Primer sequences for Illumina analysis, Table S4: Results of isotopic analysis of degas water and injected gas, Table S5: Absolute quantification of mcr-A gene via qPCR (gene copies per ml), Table S6: Table of Illumina reads for DADA2 analysis, Figure S1: Piper and Durov diagrams for groundwaters sampled at UGS Lobodice, UGS Tvrdonice, and UGS Dunajovice, Figure S2: Stability pH-Eh diagrams for carbon (a) nitrogen (b) and sulfur (c) dissolved components. Diagrams were prepared for UGS Lobodice groundwater conditions, conditions of UGS Tvrdonice and UGS Dunajovice are not significantly different. (d) Speciation of carbon dissolved species in dependence on redox potential, Figure S3: Rarefaction curve, Figure S4: Physical-chemistry well water properties in different sites and composition of archaeal community, Figure S5: The Alpha biodiversity index (Shannon) of sampled wells, Figure S6: The Beta biodiversity index (unweighted UniFrac) of sampled wells.

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